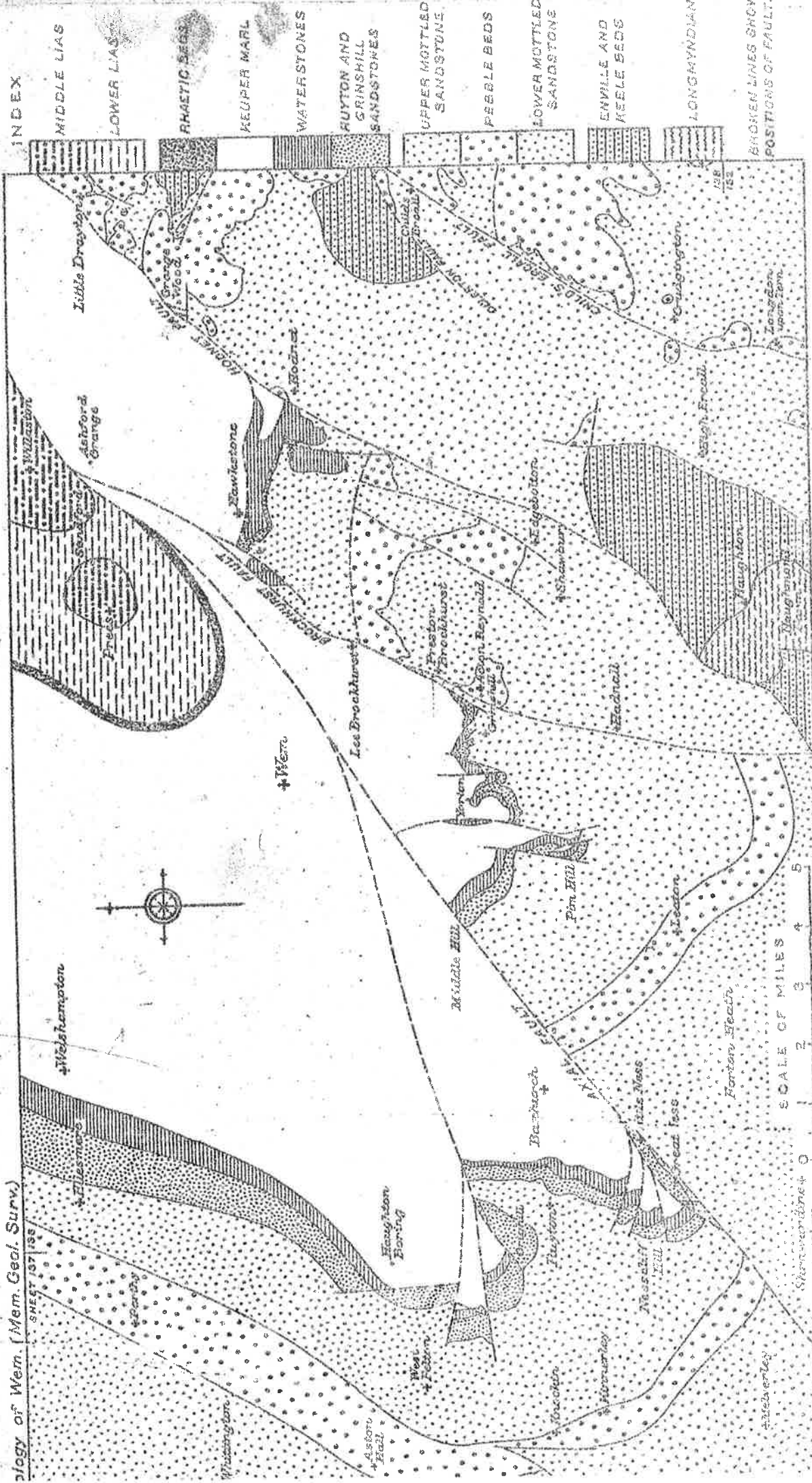


Plate II.

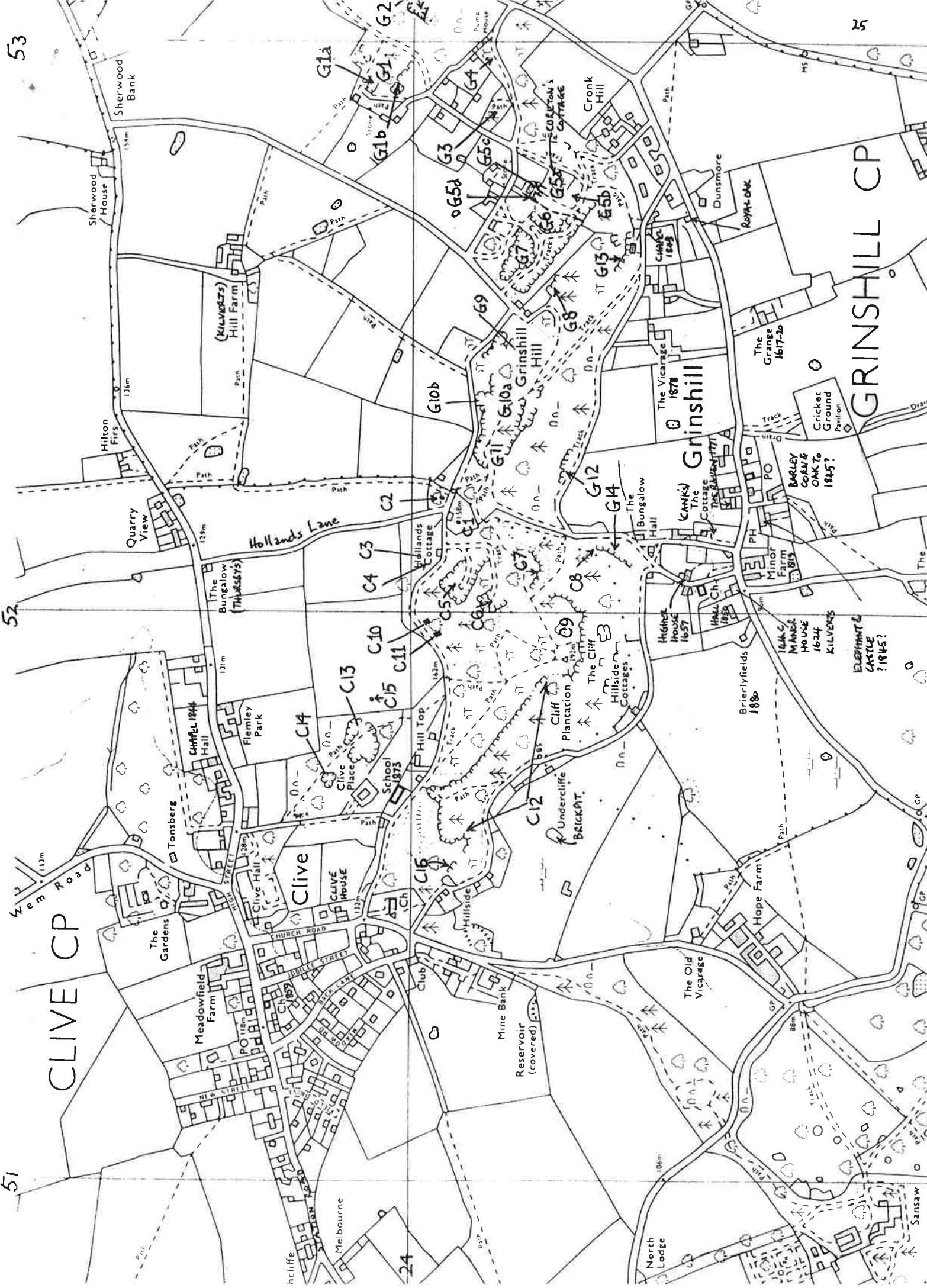


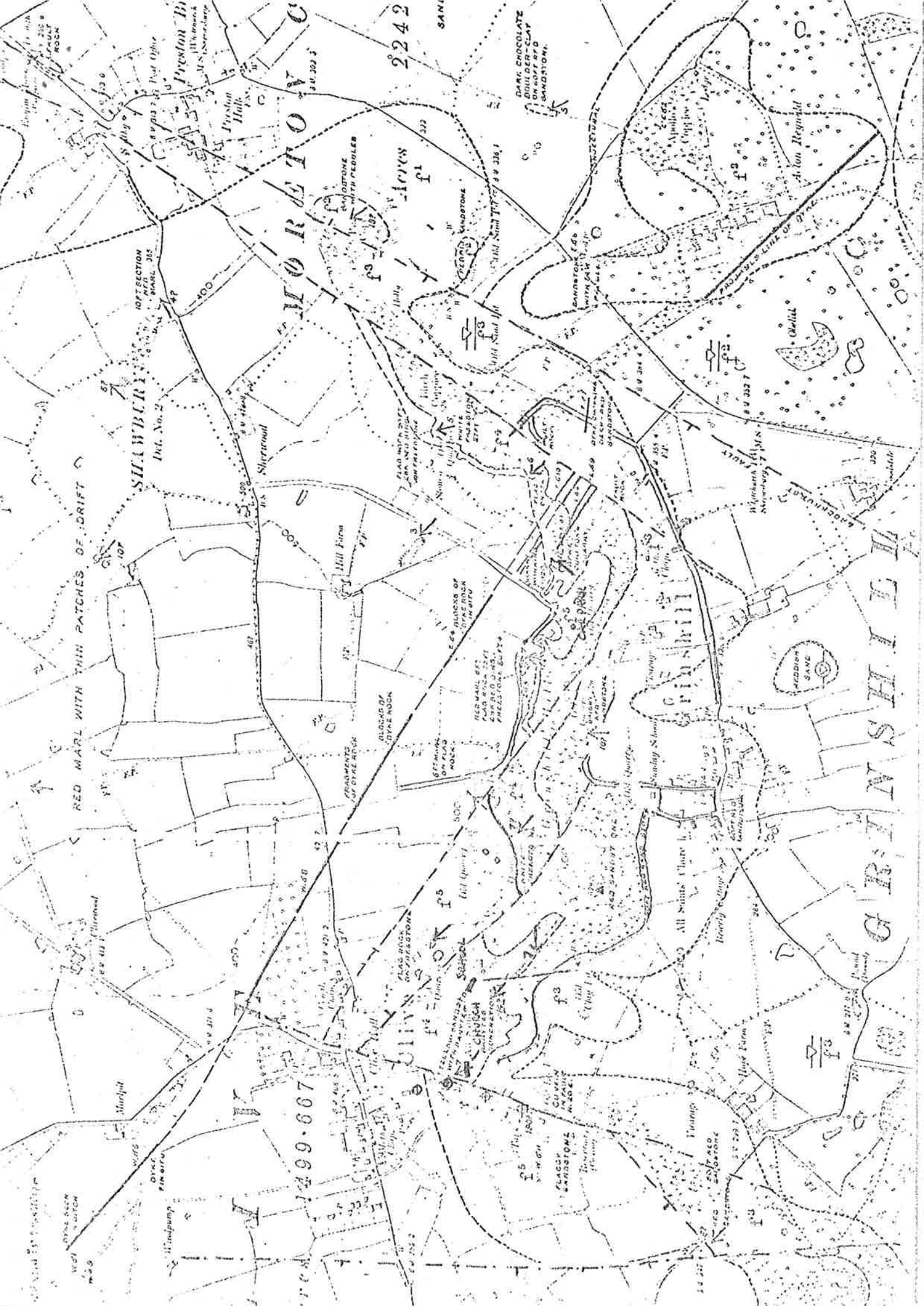
TECTONIC MAP OF THE TRIASSIC BASIN OF NORTH SHROPSHIRE

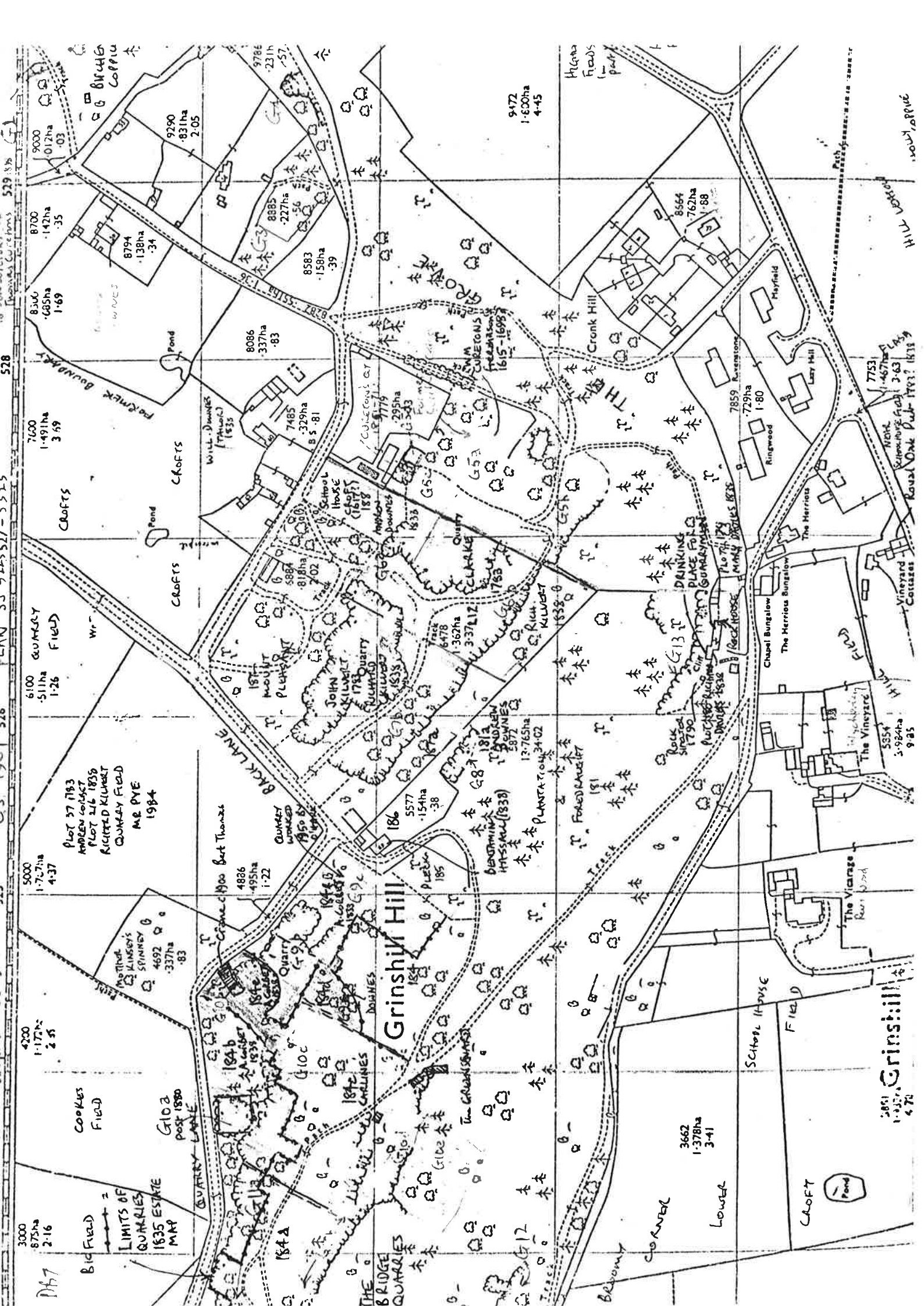
Geology of Wern. (Mem. Geol. Surv.)

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SHEET 137/138







187

Big Field

COOKES FIELD

LIMITS OF QUARRIES 1835 ESTATE MAP

GI03 post 1880 QUARRY LANE

4200 1-17ha 2-35

5000 1-72ha 4-37

Plot 37 1183 ANDREW COLCOT PLOT 216 1839 RICHARD KILBERT QUARRY FIELD MR PYE 1984

6100 QUARRY 5-11ha 1-26 FIELD

7000 1-49ha 3-59

8300 1-42ha 3-35

9000 1-01ha 0-03

9785 2-31ha 5-7

9290 831ha 2-05

8794 138ha 3-34

8300 685ha 1-69

8885 227ha 5-56

8583 158ha 3-39

8086 337ha 8-83

7485 329ha 8-5-81

5884 818ha 2-02

1874 1874

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Grinshill Hill

THE BRIDGE QUARRIES

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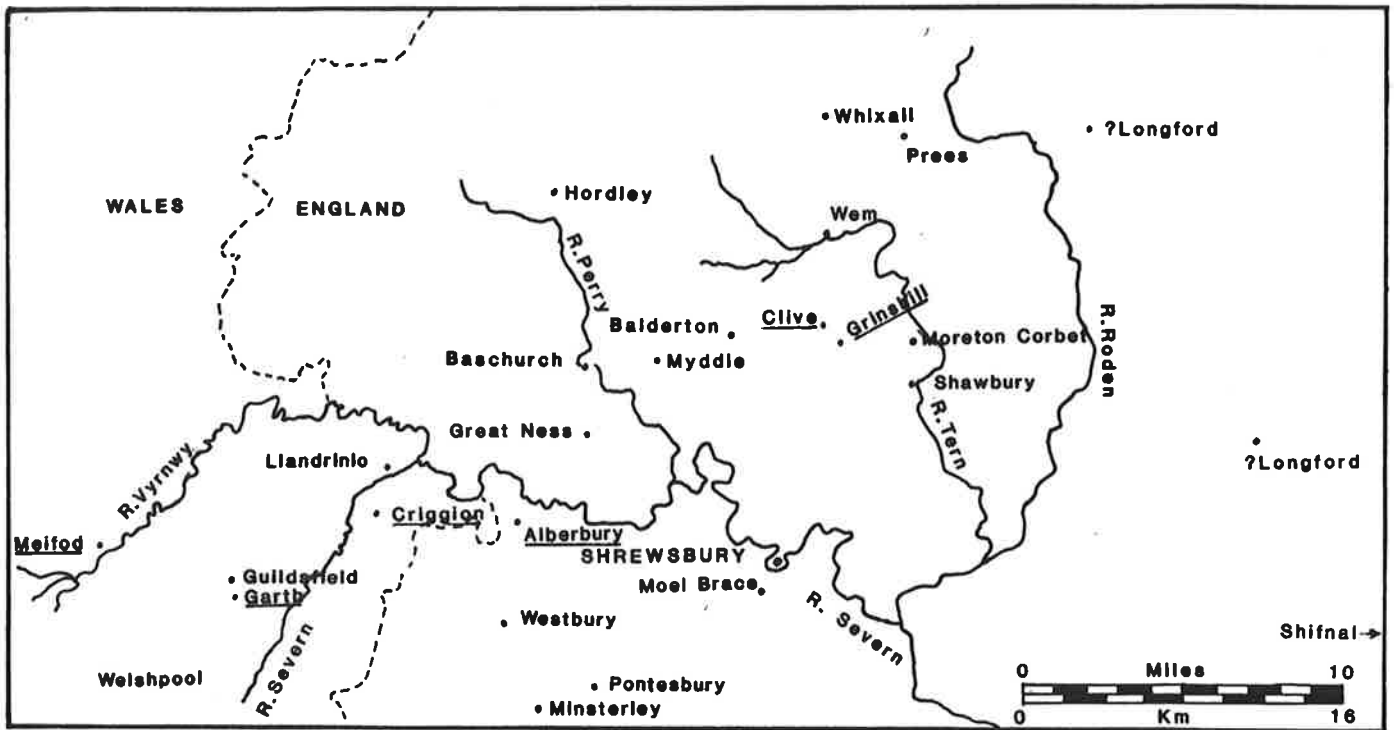
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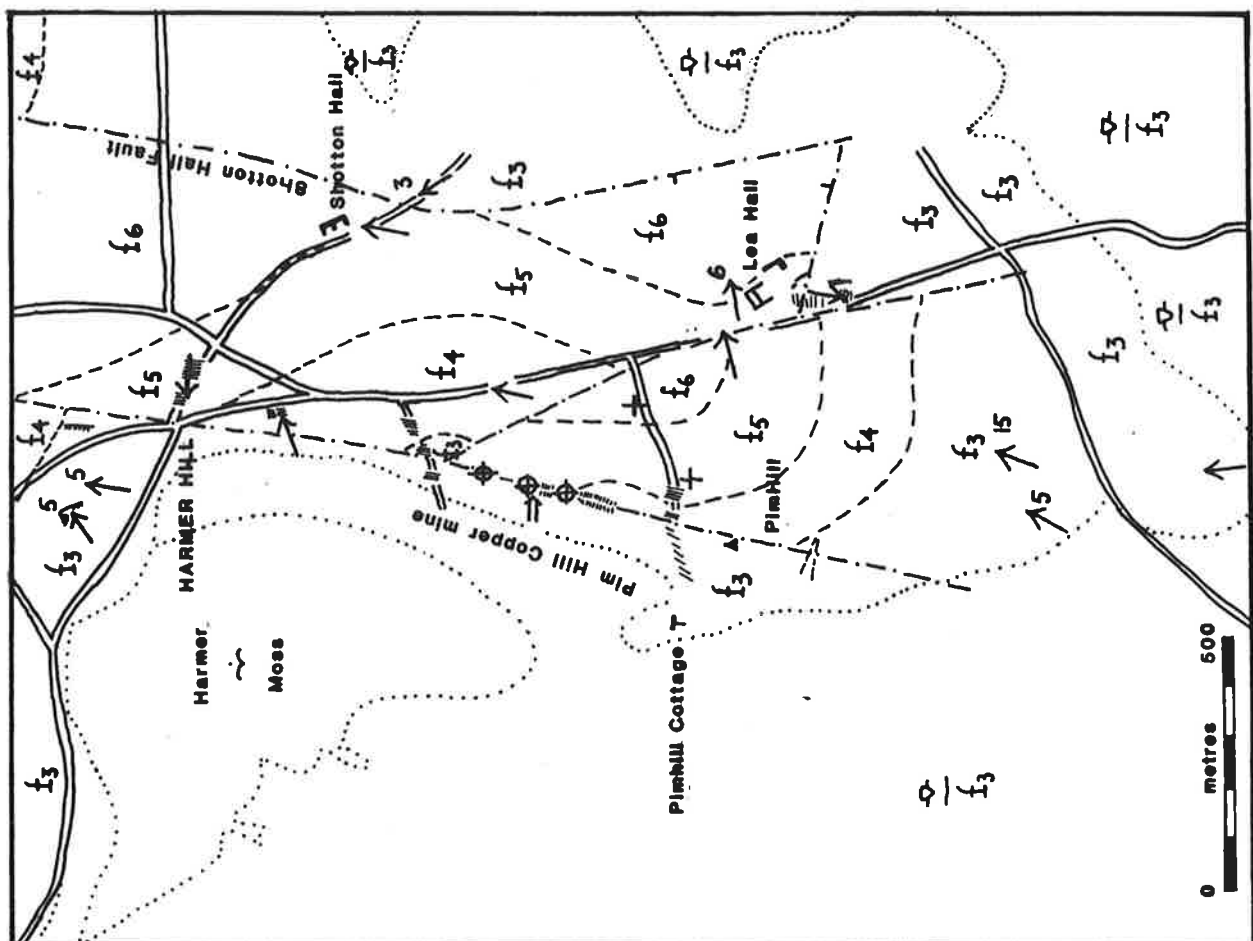
THE ROCK SUCCESSION AROUND GRINSHILL (SALOP).

Group	Visual Log	Divisions	Lithologies	Fossils	Environment	THE DYKES
SCYTHIAN		NORTHWICH ITALITE FORMATION UNAMED MUDSTONE FORMATION	Red & green interbedded siltstone and mudstones. NaCl pseudo morphs, sun-cracks, ripple marks. Thin sandstones (sheries) more common at base. Adhesion ripple bedding in thin siltstones.	Fragmentary plants - (unidentified). Trace fossils - burrows.	TRANQUIL & LOWER LOW FLOW REGIME DEPOSITS ? LAKE & LAKE MARGINAL ? INTERTIDAL (MID-TO UPPER FLATS) SURFACES SUBJECT TO AEROLIAN PROCESSES ALTERATIONS OF LATERAL & VERTICAL ACCRETION DEPOSITS LOWER LOW FLOW REGIME ACTIVATES WITH DEPOSITION FROM SUSPENSION. QUIET ? FLUVIAL REGIME ? LAKE DEPOSITS INTERTIDAL DEPOSITS (LOWER SAND FLATS)	Discovered by Murchison (1835, 1839). RC Action Reg. 2012. Extended distribution. Geol. Survey of 1920. (Pocock & Wray 1925). In two places the dyke trend is shifted by N-S faults. Dykes cut through NSF, HSF, TSE, MME. Some and in NSF, MME. Description by Dr Flett. A decomposed porphyritic dolerite, fine grained, rounded, to margin, with small porphyritic feldspar in the microcrystalline ground mass. Marginal area is vesicular and always much decomposed. (Cavities) portions are dark grey, with white feldspar phenocrysts up to 5mm. Rock is well preserved: pyroxene entirely replaced by secondary products, not possible to discern whether olivine present originally. Feldspar (labradorite) with complex twinning and slight zoned structure. Albitized in places. Feldspar in groundmass is labradorite. Magnetite, ilmenite frequent: lausene, apatite, pyrite, common chlorite, kaolinite, calcite, quartz (secondary) abundant. MAY HAVE BEEN A THOLEIITIC DOLERITE. REGARDED IN 1930S AS PROBABLY TERTIARY IN AGE AND ASSOCIATED WITH BRITO-ICELANDIC (THULEAN) IGNEOUS PROVINCE. ISOTOPIC AGE DATE 51-52 MY. (Eocene) (FITCH & MILLER) (1960S). ? THE DATE OF ALTERATION MINERALIZATION NOW THOUGHT TO BE ASSOCIATED WITH THE OPENING OF THE NORTHERN MOST ATLANTIC BY PLATE TECTONIC SEPARATION OF GREENLAND FROM EURASIA.
ANISIAN		3-13 m mean 6.5 m TRANQUIL & LOWER LOW FLOW REGIME FORMATION	Grey, light yellow, or green or pink fine grained flat bedded, thin bedded ss's interbedded with. Very thin green micaceous siltstone - mudstone laminated. Solid structures (prod, barve, flute marks; bed & flame casts); clay galls (unidentified); flat (lamination with current lamination); cross lamination; asymmetric ripple marks of transverse straight and lenticular types. Symmetric ripple marks; flat topped ripple marks; halite pseudomorphs; wrinkle marks; ripple marks on crests of ripples. Mud laminated druse structures. Mud cracks. Pinkish cavities after gypsum - anhydrite. Yellow grey sand and sst. incoherent, with some R. ridges patches: speckled with MnO2. Sometimes baryte cement.	Bones of Rhynchonella articeps Owen (Ward 1838). Suite of trace fossils including Rhynchonella articeps Owen (Ward 1838). Mainwell = Footprint D1 of Bousley.	? SUPRATIDAL DEPOSIT ? CAVITIES AFTER EVAPORATION	fine grained, rounded, to margin, with small porphyritic feldspar in the microcrystalline ground mass. Marginal area is vesicular and always much decomposed. (Cavities) portions are dark grey, with white feldspar phenocrysts up to 5mm. Rock is well preserved: pyroxene entirely replaced by secondary products, not possible to discern whether olivine present originally. Feldspar (labradorite) with complex twinning and slight zoned structure. Albitized in places. Feldspar in groundmass is labradorite. Magnetite, ilmenite frequent: lausene, apatite, pyrite, common chlorite, kaolinite, calcite, quartz (secondary) abundant. MAY HAVE BEEN A THOLEIITIC DOLERITE. REGARDED IN 1930S AS PROBABLY TERTIARY IN AGE AND ASSOCIATED WITH BRITO-ICELANDIC (THULEAN) IGNEOUS PROVINCE. ISOTOPIC AGE DATE 51-52 MY. (Eocene) (FITCH & MILLER) (1960S). ? THE DATE OF ALTERATION MINERALIZATION NOW THOUGHT TO BE ASSOCIATED WITH THE OPENING OF THE NORTHERN MOST ATLANTIC BY PLATE TECTONIC SEPARATION OF GREENLAND FROM EURASIA.
SCYTHIAN		'ESK' BED 15cm - 45m RUYTON & GRINSHILL SANDSTONE MEMBER 16m - 45m PASSAGE BEDS 5m	Yellow, white, lighter red, generally dull red. Massive, thick beds of even grain - hence a free stone. Hard on soft, hardening on exposure. Some v. large sets of cross bedding with some troughs and some signs of reactivation. In east Silicified joints: nodular barytes in the western parts near Clive. One clay horizon for a short distance. Palaeocurrent to WNW. Basal part cross-bedded at Glat near Clive School. Red and yellow greenstones	? Bones of Rhynchonella articeps Owen a Rhynchonella articeps Owen of the Turf Lizard (Sphenodon) in New Zealand. See papers of A.D. Walker 1969 Geol. Mag. 1970 G.T.S.	? FLUVIAL (IGS) DOMINANTLY LATERAL ACCRETION DEPOSITS OR MOST LIKELY AEROLIAN (THOMPSON 1970) PALAEOCURRENT FROM EAST = TRADE WIND (PALAEO MAGNETISM SHOWS 19°N OF EQUATOR ELSEWHERE AT THIS HORIZON)	fine grained, rounded, to margin, with small porphyritic feldspar in the microcrystalline ground mass. Marginal area is vesicular and always much decomposed. (Cavities) portions are dark grey, with white feldspar phenocrysts up to 5mm. Rock is well preserved: pyroxene entirely replaced by secondary products, not possible to discern whether olivine present originally. Feldspar (labradorite) with complex twinning and slight zoned structure. Albitized in places. Feldspar in groundmass is labradorite. Magnetite, ilmenite frequent: lausene, apatite, pyrite, common chlorite, kaolinite, calcite, quartz (secondary) abundant. MAY HAVE BEEN A THOLEIITIC DOLERITE. REGARDED IN 1930S AS PROBABLY TERTIARY IN AGE AND ASSOCIATED WITH BRITO-ICELANDIC (THULEAN) IGNEOUS PROVINCE. ISOTOPIC AGE DATE 51-52 MY. (Eocene) (FITCH & MILLER) (1960S). ? THE DATE OF ALTERATION MINERALIZATION NOW THOUGHT TO BE ASSOCIATED WITH THE OPENING OF THE NORTHERN MOST ATLANTIC BY PLATE TECTONIC SEPARATION OF GREENLAND FROM EURASIA.
SCYTHIAN		PASSAGE BEDS 5m	Bright, deep red, soft fine-grained x-bedded ss's. Palaeocurrent to WNW.		? FLUVIAL BRAIDED RIVER FROM SE LATERAL ACCRETION DEPOSITS ? AEROLIAN EPISODES	fine grained, rounded, to margin, with small porphyritic feldspar in the microcrystalline ground mass. Marginal area is vesicular and always much decomposed. (Cavities) portions are dark grey, with white feldspar phenocrysts up to 5mm. Rock is well preserved: pyroxene entirely replaced by secondary products, not possible to discern whether olivine present originally. Feldspar (labradorite) with complex twinning and slight zoned structure. Albitized in places. Feldspar in groundmass is labradorite. Magnetite, ilmenite frequent: lausene, apatite, pyrite, common chlorite, kaolinite, calcite, quartz (secondary) abundant. MAY HAVE BEEN A THOLEIITIC DOLERITE. REGARDED IN 1930S AS PROBABLY TERTIARY IN AGE AND ASSOCIATED WITH BRITO-ICELANDIC (THULEAN) IGNEOUS PROVINCE. ISOTOPIC AGE DATE 51-52 MY. (Eocene) (FITCH & MILLER) (1960S). ? THE DATE OF ALTERATION MINERALIZATION NOW THOUGHT TO BE ASSOCIATED WITH THE OPENING OF THE NORTHERN MOST ATLANTIC BY PLATE TECTONIC SEPARATION OF GREENLAND FROM EURASIA.

CHARTER
PERGAMON
FOUNDED
AT PERGAMON



Moreton moved Querns underground



THE GEOLOGY OF THE GRINSHILL AREA (see Fig 1 and the succession of rocks in this area)

The Grinshill area lies in the southern part of the South Lancashire, Cheshire, North Shropshire Basin (Fig.1). This basin was initiated after the deposition, uplift and erosion of the Pennine Carboniferous Basin somewhere between 300 and 270 million years ago. The North American-Eurasian plate, recently enlarged by the collision of the African plate and subject to east-west wrenching, became relatively stable and affected by east-west tension, so that elongate basins (or half-rifts) developed. Deposition of the Sherwood Sandstone Group and the later Mercia Mudstone Group took place over a period 270-213 million years, sedimentation roughly keeping pace with subsidence. Some of the details of the rock succession in the local area are given in diagrammatic form. The area was folded and faulted at a much later date, the earth movements being sustained episodically over long periods. The area was intruded by basic magma 250 million years ago as the North Atlantic Ocean opened, was mineralised, and was covered (in the period 24000-13000 yrs) by the last advance of the last ice sheet (the Devensian ice).

The minerals and their distribution

The Triassic rocks within the Basin are essentially silica rich sediments cemented by silica (SiO_2), calcite (CaCO_3) and iron oxides/hydroxides, the latter imparting the characteristic red colouration. A considerable number of other minerals occur in the rocks including feldspars and clay minerals, but the ones of interest in the mining areas were compounds containing copper, lead, cobalt or silver. All of these elements were mined at one time or another in the region.

The principal minerals are malachite, azurite and chrysocolla (copper); galena, pyromorphite and cerussite (lead); asbolite (cobalt, nickel and manganese); wad (manganese) and barite (barium sulphate). Occurring with these metals are arsenic, silver, vanadium, and molybdenum. A list of minerals so far recorded from the region is presented in Table 1.

TABLE 1. Minerals recorded or reported from the metalliferous deposits of the Cheshire-Shropshire Basin. Dominant species * (As known in 1979).

COPPER	Native Copper	LEAD *	*Galena	IRON	Pyrite
	Chalcocite		Massicotite		Haematite
	Chalcopyrite		Minium		Goethite
	Bornite		*Cerussite		Limonite
	Enargite	MANGANESE	*Pyromorphite		Turgite
	Cuprite		Arsenical pyromorphite		Pisanite
	*Malachite		Anglesite		Scorodite
	*Azurite				Beudantite
	Chrysocolla	COBALT		ZINC	Sphalerite
	Tyrolite		Pyrolusite		
	Planchite		*Wad		
	Olivinite			MOLYBDENUM	
	Brochantite	VANADIUM	*Asbolite		Wulfenite
	Libethanite		Erythrite	BARIUM	
	Liroconite		Bravoite		*Barite
	Caledonite		Coballite 42		Witherite
	Langite		Roscoelite	NICKEL	
	Covellite 42		Vanadinite		Bravoite 42
	Djurleite		Mottramite	SILVER	Gersdorffite 42
					Tetrahedrite

The metallic minerals occur with a barite 'gangue' as disseminations within the sandy and pebbly Triassic rocks, coating sand grains, infilling pore-spaces and forming a patchy rock cement. The mineralised rock often has ill-defined contacts and varies foot by foot in metal grade, thickness and width, forming lenses, streaks and irregular patches of mineralised ground. In detail the distribution and association of the minerals is complex but essentially follows a pattern which can be summarised thus :-

1. The mineralisation is predominantly copper-cobalt-manganese-barite. Only at Alderley Edge and Mottram Mine was lead present in economic quantities.
2. Barite and wad (manganese) have a wide geographic distribution around the Basin, while copper-cobalt-(lead) minerals are more restricted to areas proximal to faults.
3. The minerals occur through a vertical range of some 6,000 feet of Triassic rock but are most abundant within the top of the Mottled Sandstone Formation and particularly in the overlying Keuper Sandstone Formation immediately below the regional 'seal' of the Waterstones. In addition Keuper Sandstone or Mottled Sandstone may be mineralised in faulted contact with Waterstones or Mottled Sandstones.

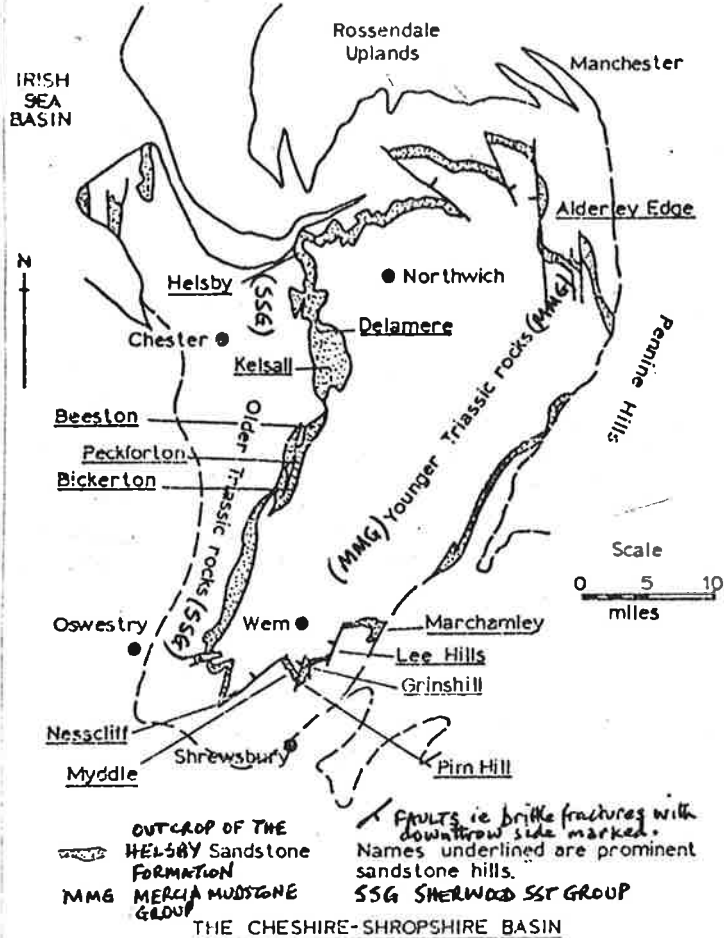


FIG 1

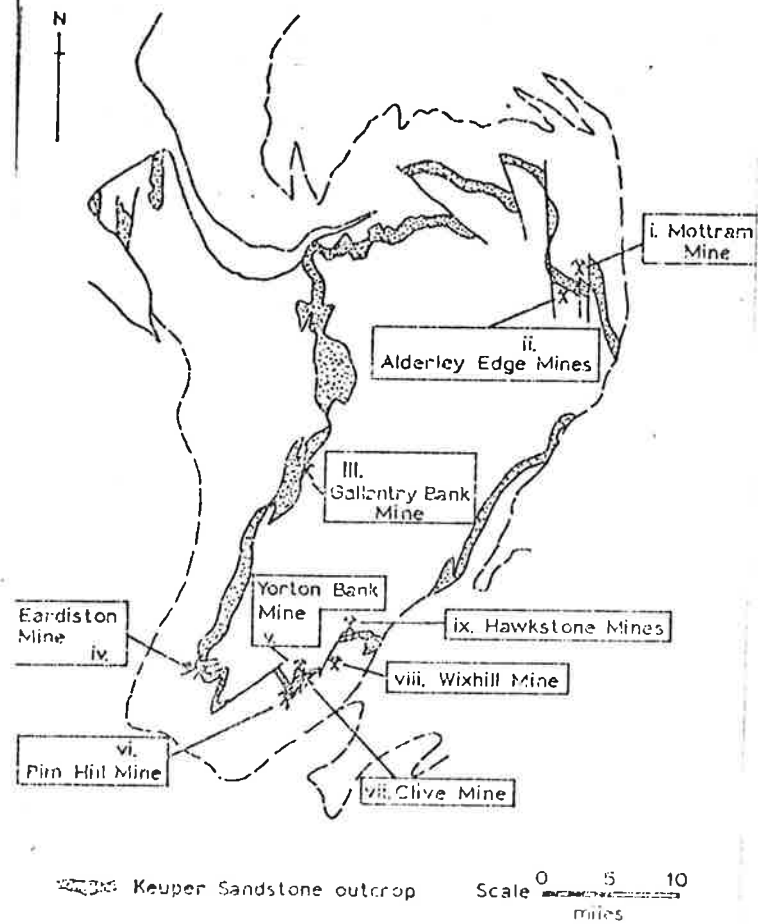
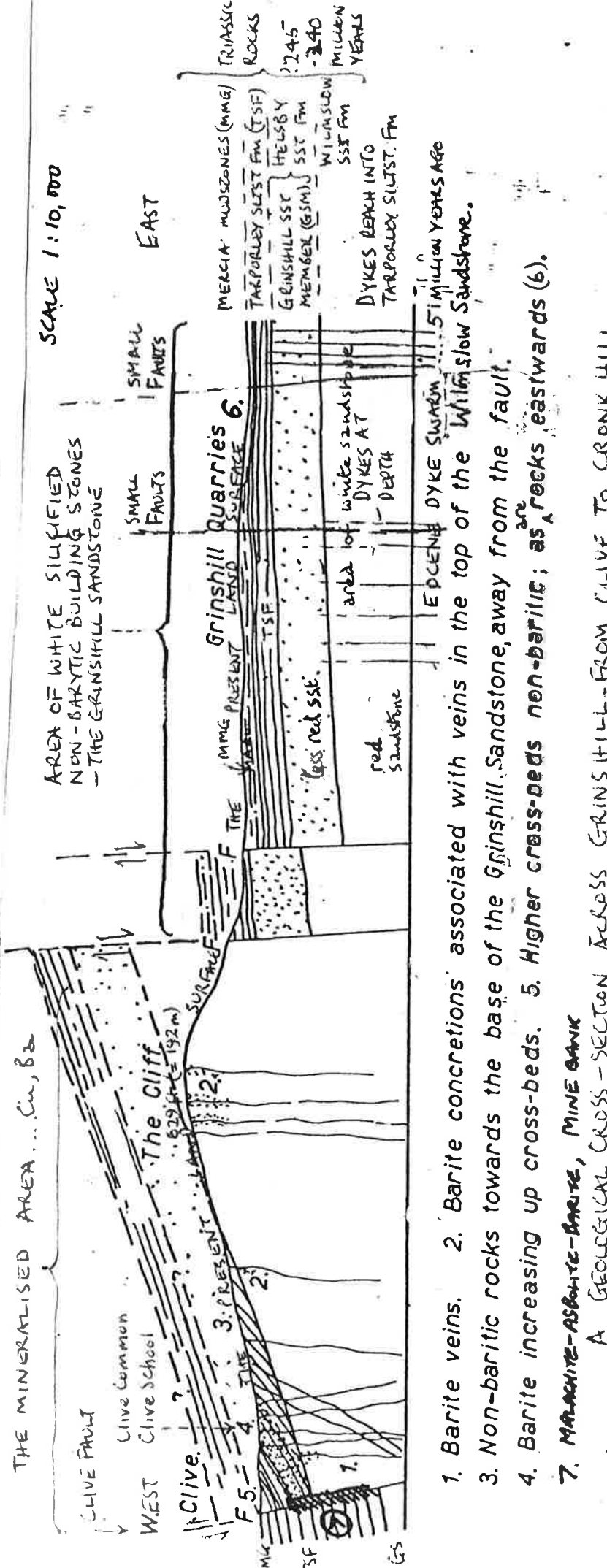


FIG 2. MINES IN THE CHESHIRE-SHROPSHIRE BASIN



The Mines

There are nine principal mining sites in the Cheshire-Shropshire Basin. All of them worked copper ores which were associated principally with cobalt at Eardiston and Clive; cobalt and silver at Bickerton and cobalt, lead and silver at Alderley Edge and Mottram St. Andrew. All sites had associated barite and wad. Little is known of the ores at Yorton, Pim Hill, Wixhill and Hawkstone, but they are likely to have been copper-cobalt in character. Vanadium was recorded at Pim Hill.

The regional setting of these mining sites is shown in Figure 2. At Alderley Edge and Mottram St. Andrew the ores were dispersed through faulted ground; at Eardiston, Bickerton, Clive, Pim Hill, Yorton and Wixhill the ores were mined from rocks adjacent to one principal fault, while at Hawkstone Park, the trial mines are in ground not adjacent to faults.

The date of first discovery of these mineral deposits is unknown, but they are likely to have attracted early attention by virtue of the brightly coloured copper compounds or the sterility of the ground around the mineral occurrence.

Mining by the Romans has been postulated for the Alderley Edge Mines (Roeder 1901)³, (Roeder and Graves 1905)⁴; for the Mottram Mine (Roeder 1901)³; and for the Clive and Hawkstone Mines (Wood in Watkin 1879)⁵. Stone artifacts discovered at Alderley Edge and Mottram Mine have led to the conclusion that mining began prior to the Roman Occupation (Boyd-Dawkins 1875)⁶.

It is likely that mining had taken place at all the sites prior to the latter half of the 1700s, with the possible exception of Eardiston Mine (Carlton In Press)². By the late 1700s the region was supporting a flourishing copper mining and smelting industry, though the most productive period of mining was undoubtedly the thirty years period 1850-1880 when the mines at Mottram, Alderley Edge, Bickerton, Eardiston, Clive, Pim Hill and Wixhill were worked. West and Wood Mines, Alderley Edge were worked intermittently in the early years of the 1900s and leases were sought or obtained for the Bickerton and Eardiston Mines. Various individuals and companies have shown a passing interest in the Alderley Edge, Bickerton, Eardiston and Clive Mines in recent years but it is most unlikely that any of them will be reopened.

All the mines seem to have been worked by initial open pitting and trenching, followed by shaft or adit access to underground mining where the ore was removed in the normal way by picking or later using drilled shot holes and gunpowder. The stopes produced were either linear or irregular, depending upon the form and nature of the orebody. In most of the mines timber props were used sparingly or not at all, the rocks standing without the need for artificial support.

Only the Alderley Edge Mines appear to have been dry workings, although the Hawkstone trial mines were probably dry also. All the other mines were wet, some alarmingly so, and required the use of steam power to dewater them. Clive, Bickerton and Eardiston all had, at sometime or other, engine houses and steam engines, and some form of motive power may have been used at Yorton and Wixhill for pumping and hoisting.

The copper ores were initially treated on site or nearby by smelting, however by the early 1800s the ores were too low grade to be treated economically by direct smelting. In 1857 William Henderson, company chemist for the Alderley Edge Mining Company, patented a process for copper extraction from poor copper ores using an acid leaching technique (Henderson 1857)⁷. This method was later adopted at Bickerton, Eardiston and Clive.

The following review of the Cheshire-Shropshire Basin mines is based on archival, field and verbal researches and is intended to shed new light on these old mines. For each mining site the location is given with the National Grid Reference. The present (1980) surface features and remains are described and a brief outline of the mining history, geology and related features given. References for each locality are given, where these are generally available.

Commencing in the north-east corner of the basin, the mines are described in anticlockwise order around the region, as numbered in Figure 2.

vii. CLIVE (GRINSHILL) MINE

Location: On the south and south-east sides of Clive village, north Shropshire. The workings extend from NGR SJ5150.2425 to SJ5130.2363.

Geology: The Clive Fault runs under Clive village throwing Marls and Waterstones to the west into contact with Grinshill Sandstone (Keuper Sandstone) and Mottled Sandstone to the east, the two latter units being mineralised. The fault dips steeply westwards, being almost vertical. The ore was found in rocks leached of their iron oxide-hydroxide content, this material forming an 'iron-pan' adjacent to the main orebody. The ore was a mixture of malachite with azurite and asbolite in a baritic sandstone adjacent to the Clive Fault. The mineral impregnation pinched and swelled, the average stopping width in the 1860s being 15-20 feet, and was tested over a lateral distance of some 600 yards to a maximum depth of 240 feet. The ore grade and volume decreased dramatically below 180 feet. The granular, grey-white host rock was in places speckled with abundant asbolite (cobaltian wad) an occurrence termed 'pepper and salt rock'. In 1862 ore raised averaged 3½% copper.

Mining: At least six shafts, numerous surface excavations and an adit were sunk and driven to test and exploit the orebody at Clive. Ore was stoped and trammed to the hoisting shaft, which in the early years of mining was at the north end of the sett. Worked to a depth of 240 feet, the levels below 180 feet were used for haulage only. A 6 feet by 9 feet shaft was sunk to a depth of approximately 160 feet at SJ5139.2386 in the hangingwall of the deposit. Originally a hoisting shaft adjacent to the mine plant area of the 1860s, this shaft was deepened in the early 1900s and is still used as a source of water, being equipped with pumps.

History: Clive Mine is yet another working in part attributed to the Romans. First mooted by Samuel Wood in 1879 (in Watkin 1879)⁵, the evidence is very shaky and hinges on the similarity between certain levels at Clive and those of proven Roman origin on Llanymynech Hill near Oswestry. As at Pim Hill, Wixhill and Hawkstone Park, the presence of copper at Clive was well known by the late 1600s. A considerable portion of the Clive Mine may have been excavated in the 1600s, and 17th century tools and artifacts have been discovered in the mine.

The mine lease was held by John Payne of Droitwich in 1739, and over the next 120 years several interested individuals and companies sought to reopen and mine the Clive orebody.

A prosperous episode of mining activity took place between 1862-1869. During this period the Clive Mining Company came into being after obtaining a lease to work the orebody from the landowner, Mr. Gardner. Some 30-40 men were employed at the mine, some of them Cornishmen and others experienced copper-sandstone miners who had worked in the Alderly Edge Mines.

They greatly extended Clive Mines; ore was raised, crushed between cogs and acid leached in 24 large stone tanks. Copper was precipitated from the solution in wooden troughs using scrap iron.

The Clive Mining Company, initially under the care of mine manager J. T. Harris who also briefly took charge of the Rednal (Eardiston) Mine about this time (Carlton in press)², constructed workshops and an engine house at the Clive Mine. The present Mine Cottage at Clive is one of the original mine buildings. On closure in 1869 the stone leaching tanks were sold to the Bryntail Mine near Llanidloes, Powys, Central Wales. Here they were used in the dressing of barytes. Restored some years ago by the Department of the Environment, the Bryntail Mine still possesses 8 of the Clive tanks, which can be seen, in various stages of restoration at NGR SN915869. The site is described by Bick (1977)¹⁹.

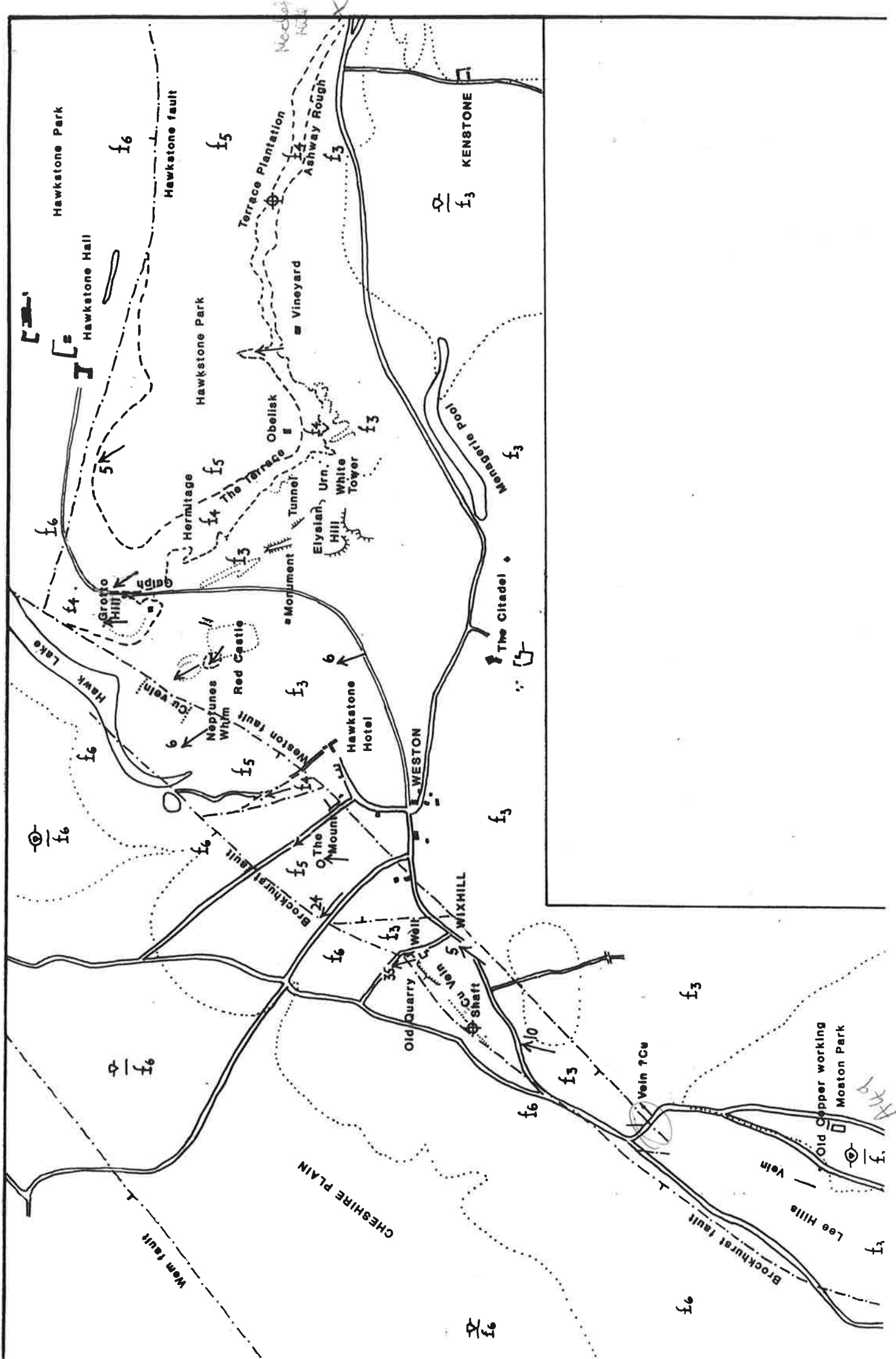
Output: The total output of ore from the Clive Mine is unknown.

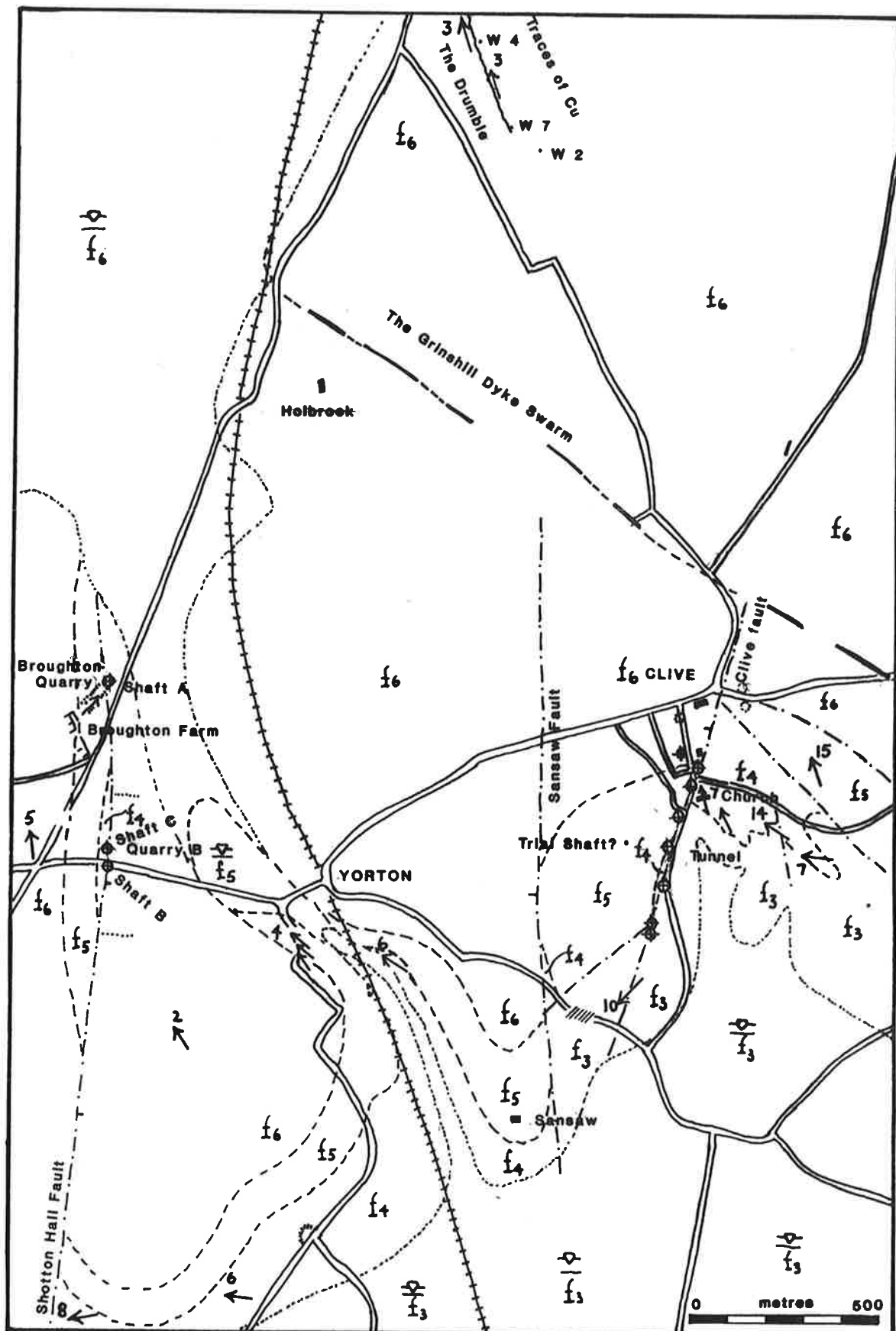
Surface: There is little to see on site at Clive. The best tangible mining relics are at Bryntail Mine some 60 miles away! The deepened hoisting shaft is still accessible, within a building adjacent to Mine Cottage on private property. Due south of Mine Cottage in a plantation a fence marks the position of a 'shaft that appeared in the 1950s due to subsidence of its infill, and 200 feet further south are two recent 'holes' where the ore was stoped to surface.

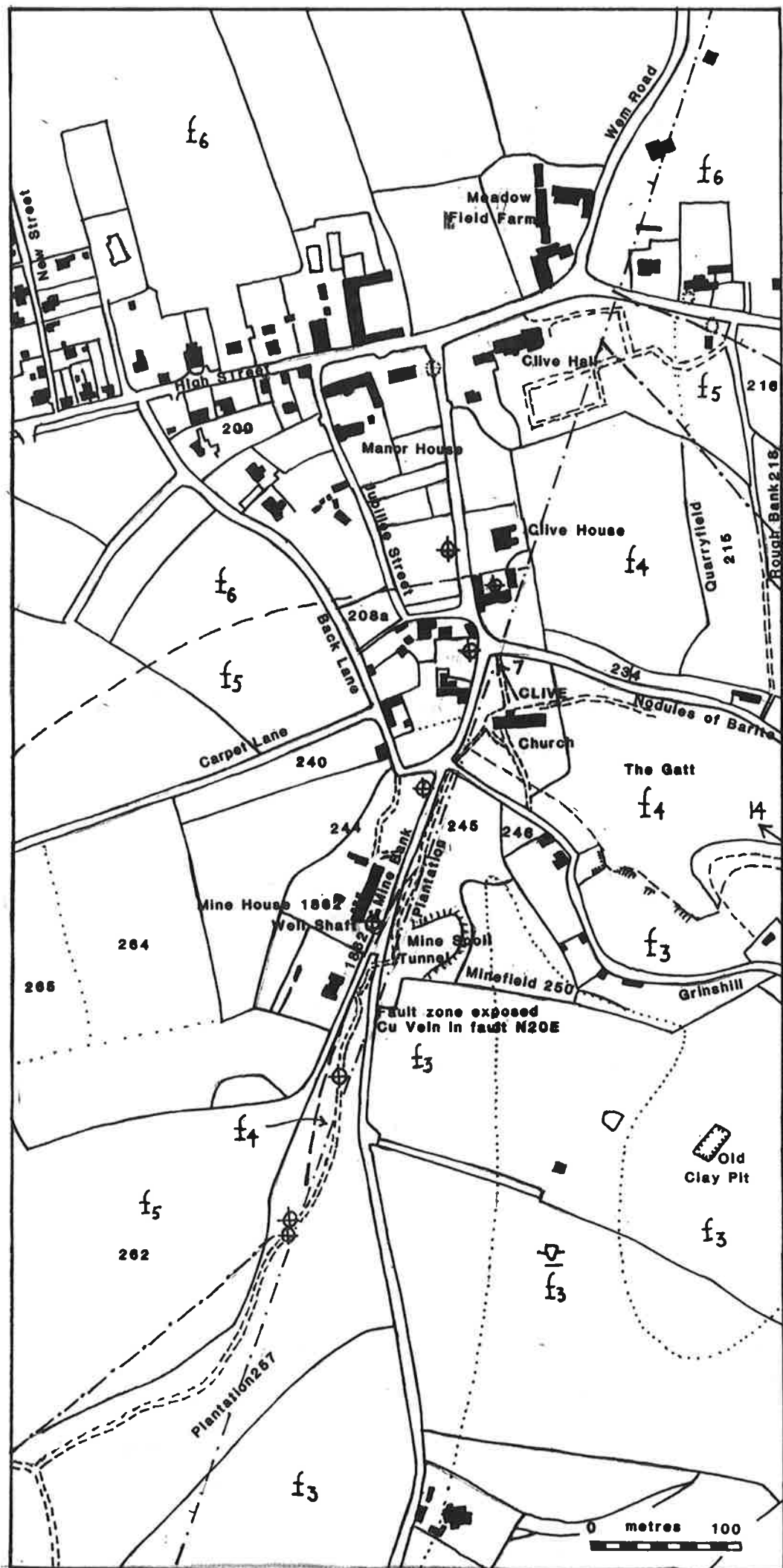
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Pocock and Wray (1925)¹⁶ pages 90 and 99.
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GEOCHEMISTRY OF THE GRINSHILL DYKE SWARM D.B. THOMPSON & J.A. WINCHESTER (UNIV. KEELE)

BL MORNING

Sample CODE	8509	8510	8511	8512	8513	8514	8515	8516	8517
	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke
S102	51.68	53	53	50.85	53.01	52.1	52.39	52.73	53.36
T102	3.57	2.8	3.08	3	3.5	3.69	3.15	3.01	2.84
AL203	29.66	24.71	26.8	23.56	24.37	28.15	27.62	24.31	20.17
FE203	.98	5.16	1.95	8.62	4.91	1.43	2.55	5.13	8.92
FeO	0	0	0	0	0	0	0	0	0
MNO	.05	.08	.06	.05	.05	.06	.1	.05	.11
MGO	.51	1.29	.54	1.87	.92	.39	.6	1.23	1.81
CaO	1.02	.99	.9	.95	.92	.96	.92	1.45	1.7
Na2O	.13	.11	.17	.09	.15	.19	.14	.17	1.27
K2O	4.06	3.68	3.54	1.93	4.8	4.21	4.42	3.51	2.67
P2O5	.53	.49	.49	.46	.46	.5	.46	.39	.34
H2O+	8.73	7.88	7.41	9.09	7.64	8.53	8.54	8.88	6.63
S	.02	.01	.02	.02	.02	.02	.02	.01	.01
BA	317	412	493	745	355	361	398	322	929
CE	86	77	114	81	112	100	91	52	55
CL	387	59	90	348	1331	863	1286	4306	136
CR	54	18	39	26	45	77	60	52	11
CU	1	36	17	47	43	20	19	19	29
GA	16	24	18	21	20	17	20	18	23
LA	51	41	57	49	24	44	43	30	11
NB	15	12	15	12	13	14	10	9	10
ND	58	43	64	40	76	54	66	76	43
NI	11	22	12	61	20	8	21	23	60
PB	59	129	101	222	194	116	133	134	59
RB	41	70	62	47	60	44	55	73	35
SR	38	66	54	65	75	44	41	50	81
TH	5	6	9	0	5	4	2	3	3
V	406	345	392	490	498	546	464	439	420
Y	92	96	109	103	104	99	85	81	63
ZN	83	150	97	213	166	85	81	96	40
ZR	376	290	337	253	307	319	265	249	238

CHARES

Sample CODE	8518	8519	8520	8521	8522	8523	8524	8525	8526
	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke	1 dyke dyke dyke
S102	49.42	52.89	45.97	44.34	45.99	47.23	53.04	56.71	56.81
T102	2.32	3.24	2.79	2.63	2.4	2.18	2.65	2.68	2.73
AL203	17.09	21.09	19.41	18.52	16.22	14.46	21.66	18.84	20.06
FE203	9.4	5.11	7.46	6.64	8.7	9.33	7.96	7.24	6.01
FeO	0	0	0	0	0	0	0	0	0
MNO	.08	.05	.11	.08	.1	.11	.16	.09	.08
MGO	2.26	1.99	1.66	2.24	3.54	3.42	.94	1.61	1.27
CaO	7.72	2.6	6.53	8.92	8.78	9.12	4.43	.83	.87
Na2O	2.21	.33	.61	.27	.62	1.5	3.02	.27	.19
K2O	1.53	3.88	2.93	2.07	1.63	1.15	2.28	4.48	4.97
P2O5	.28	.41	.38	.36	.39	.41	.78	.63	.73
H2O+	7.46	8.19	10.55	12.97	9.87	9.58	4	6.34	6.5
S	.03	.01	.01	.02	.11	.07	.02	.01	.01
BA	3329	135	288	1719	512	1739	1497	659	503
CE	50	71	51	55	53	23	59	135	118
CL	324	96	67	73	59	69	94	67	74
CR	1	50	8	4	1	1	1	1	1
CU	62	90	100	94	83	74	35	31	16
GA	20	28	29	23	26	20	21	20	17
LA	8	26	17	22	5	10	19	38	33
NB	10	14	12	12	15	14	12	19	21
ND	37	34	44	38	44	23	18	64	82
NI	37	32	49	35	11	37	37	14	26
PB	27	2342	1077	1521	662	107	78	65	78
RB	24	49	50	28	24	20	27	64	61
SR	189	34	45	87	80	123	228	35	41
TH	3	25	9	12	6	4	2	5	9
V	380	501	409	407	344	273	436	284	242
Y	54	85	87	76	75	81	66	108	127
ZN	34	5424	9933	5627	4823	1617	42	27	21
ZR	203	296	263	258	261	279	232	409	461